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Advancing Interpretable AI/ML Methods for Deeper Insights and Mechanistic Understanding in Earth Sciences: Beyond Predictive Capabilities

Kev Points:

- · The fractal dimension calculated using Shen & Li's model optimally quantifies the pore structure of tight sandstone
- The fractal dimension is the most important factor for permeability prediction
- The explainable hybrid model outperforms traditional models, providing insights into complex and hidden feature interaction predictions

Supporting Information:

Supporting Information may be found in the online version of this article.

Correspondence to:

F. Jiang, jfjhtb@163.com

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Author Contributions:

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Project administration: Fujie Jiang,

Zezhang Song

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A Novel Explainable Hybrid Model for Permeability **Prediction of Tight Sandstone Using PI-DeepFM Machine Learning Algorithm**

Lina Huo^{1,2}, Fujie Jiang^{1,2}, Liu Cao^{1,2}, Di Chen^{1,2}, Renda Huang³, Xiaowei Zheng^{1,2}, Guangjie Zhao^{1,2}, Weiye Chen^{1,2}, Zezhang Song^{1,2}, Libin Zhao⁴, and Yuanyuan He⁴

¹National Key Laboratory of Petroleum Resources and Engineering, China University of Petroleum (Beijing), Beijing, China, ²College of Geosciences, China University of Petroleum (Beijing), Beijing, China, ³Institute of Energy, Peking University, Beijing, China, ⁴Taxinan Exploration and Development Company, PetroChina Tarim Oilfield Company, Kolar,

Abstract Permeability is essential in controlling hydrocarbon fluid distribution and production in tight sandstone reservoirs. However, most conventional permeability prediction models are low-accuracy and cannot evaluate the importance of factors quantitatively because of the complex interplay among them, stemming from the complex pore structure of tight sandstone. Although machine learning (ML) constitutes an advanced solution, existing ML models often neglect the importance of explainable feature selection and interactions, thereby significantly restricting accuracy and reliability. Here, we proposed an explainable hybrid ML model, namely, PI-DeepFM, for permeability prediction. In this model, the Permutation Importance (PI) algorithm for feature selection is incorporated with Deep-Learning Neural Network and Factorization Machine (DeepFM) algorithm for feature interaction and prediction. To establish a more accurate feature set for prediction, we conducted various experiments, identified Shen & Li's fractal model as the best method for characterizing tight sandstone pore structure, and analyzed the factors influencing permeability through geological and correlation mathematical analysis methods. The quantitative feature evaluation results of the PI-DeepFM revealed that the fractal dimension, derived from Shen & Li's fractal model, influenced permeability the most, followed by skewness, Clay, porosity, Carbonate and sorting coefficient. This indicates that fluid seepage is dominated by pore structure complexity and connectivity. The PI-DeepFM model demonstrated a remarkable reduction of 25.76% in the root mean square error compared with that of conventional methods.

1. Introduction

Tight sandstone gas, a clean and green unconventional fossil fuel, accounts for approximately 75% of the total unconventional natural gas resources, playing a critical role in meeting the growing global energy demands (Jia et al., 2022; Qiao et al., 2020; Sun et al., 2013; Worden et al., 2018). Permeability is a macroscopic parameter that describes the pore structure of porous media and governs fluid distribution and production in tight sandstone reservoirs (Behrang et al., 2016; X. Chen & Yao, 2017; Golab et al., 2010; X. Zhou et al., 2024). Permeability prediction based on its relationship with other petrophysical properties or pore structure parameters has emerged as a key focus in reservoir evaluation research, particularly for tight sandstone (R. Guo et al., 2020; Rezaee et al., 2012). Nevertheless, the heterogeneous and complex pore structure of tight sandstone, which results in various factors affecting permeability and intricate interactions among them, presents challenges in establishing an accurate permeability prediction model (Qiao et al., 2022).

Traditional permeability prediction models based on Poiseuille's equation or Darcy's law, such as the Winland and Pittman models, are usually employed to predict permeability and are commonly applied in conventional carbonate and sandstone reservoirs (Y. Guo et al., 2022). Tight sandstone exhibits a complicated pore structure and high heterogeneity, with its permeability resulting from the comprehensive influence of multiple geological factors. Therefore, a single or several feature parameters in empirical models may not fully reflect the seepage characteristics or might not be the critical factors influencing permeability. These will lead to inapplicability and large errors in prediction (Al-Kharra'a et al., 2023; R. Guo et al., 2020; Liu et al., 2018; Qiao et al., 2020). Previous studies have revealed that the main influencing factors include (a) macro petrophysical parameters like porosity (Dou et al., 2021; Z. Li et al., 2020; F. Wang & Wang, 2022); (b) experimental quantitative characterization parameters of pore structure, including high pressure mercury intrusion (HPMI) parameters and fractal dimension

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Resources: Libin Zhao, Yuanyuan He Software: Liu Cao Supervision: Fujie Jiang, Di Chen, Renda Huang, Xiaowei Zheng, Guangjie Zhao, Weiye Chen, Zezhang Song, Libin Zhao, Yuanyuan He Validation: Lina Huo, Liu Cao Visualization: Lina Huo, Liu Cao Writing – original draft: Lina Huo Writing – review & editing: Lina Huo, Di Chen, Renda Huang, Xiaowei Zheng, Guangjie Zhao

(*D*) (Dai et al., 2021; Lai et al., 2018; F. Wang et al., 2019); and (c) the mineral composition, of which different mineral types and genetic mechanisms affect the pore size, shape and distribution, in turn influencing permeability (Qiao et al., 2022; Q. Zhang et al., 2022; F. Zhu et al., 2018). Although mercury intrusion porosity (MIP)—based models have been extensively used for pore structure characterization, their optimal selection remains understudied, particularly in the Bozi-Dabei area of the Tarim Basin, which is critical for precise pore structure characterization (Dou et al., 2021; K. Li, 2010; Z. Zhang & Weller, 2014). To construct a more accurate feature set for prediction, it is essential to characterize the pore structure better, analyze the factors affecting permeability at greater depths, and quantitatively assess their importance.

Machine learning (ML) has been extensively applied in geological research due to process data efficiently and its ability to capture non-linear relationships (Elmorsy et al., 2022; Jiang et al., 2023; Ma et al., 2022; L. Wang et al., 2020, 2021). ML provides a promising solution to quantitative feature importance and prediction problems with complex feature interactions. Given the diversity and complexity of influencing factors, effective feature selection to eliminate redundancy and irrelevance is essential before predicting. The traditional feature selection methods, such as Pearson correlation (Pearson, 1901) and Principal Component Analysis (PCA) (Jolliffe, 2002), are more suitable for linear relationships, whereas the Spearman correlation (Spearman, 1904) is better suited for monotonic relationships among features and is robust to outliers. With the continuous development of ML, linear regression algorithms, such as Lasso regression (Hastie et al., 2015), achieve sparse feature selection through regularization, but they are still limited to linear modeling. Tree-based algorithms like Decision Trees (DT) (L. Breiman et al., 1984), Random Forests (RF) (L. Breiman, 2001; Strobl et al., 2008), Extreme Gradient Boost (XGBoost) (T. Chen & Guestrin, 2016), and LightGBM (Ke et al., 2017) can more effectively capture nonlinear relationships between features and target variables (Kuhn & Johnson, 2018; Strobl et al., 2008). However, these models have varying approaches and sensitivities to feature handling, necessitating the selection of feature selection methods that conform to the predictive model (Guyon et al., 2002; Kuhn & Johnson, 2018). And the regular prediction algorithms face specific challenges, such as the need for pre-training in Factorized Neural Network (FNN), limitations in capturing low-order feature interactions in Product-based Neural Network (PNN), the necessity for manual feature engineering in Wide & Deep models, and the risk of overfitting in Random Forest (RF), especially with an excessive number of trees, which can limit their ability to capture complex feature interactions effectively. In conclusion, existing ML-based permeability prediction models exhibit problems including a lack of effective feature selection, reliance on certain prior knowledge, limited interpretability, and deficiencies in mining feature interaction information, which affect their trustworthiness and accuracy (L. Breiman, 2001; Cortes & Vapnik, 1995; Mahdaviara et al., 2021; Rasmussen & Williams, 2005; Wold et al., 1987; W. Zhang et al., 2021). There is a pressing need for advanced, explainable ML algorithms that can quantify the importance of factors to eliminate redundant and irrelevant features and effectively mine intricate interactions.

Here, we propose an explainable hybrid ML model (PI-DeepFM) for permeability prediction, which incorporates the Permutation Importance (PI) algorithm with the Deep-Learning Neural Network & Factorization Machine (DeepFM) algorithm. The PI algorithm, which is well-known for its robust interpretability and suitability for structured small sample data and any model, enables feature selection via the computation of the importance score of each feature (Altmann et al., 2010; Fisher et al., 2018). The DeepFM algorithm can be used to simultaneously mine linear, pairwise, and high-order feature interactions without manual feature engineering (H. Guo et al., 2017). It represents the cutting-edge in recommendation systems and has demonstrated exceptional performance across domains such as information retrieval, agriculture, and meteorology (Khaki et al., 2020; Yi et al., 2022). To our knowledge, this study presents the first application of this algorithm in the geological field. In this study, we applied the hybrid ML model to the Cretaceous Bashijiqike Formation (BS Fm.) in the Bozi-Dabei area, Tarim Basin. We conducted varied experimentsto characterize the pore structure, identified the optimal model from six fractal models and analyzed the factors influencing the tight sandstone permeability by correlation analysis with geological knowledge to construct a more accurate feature set. The feature set was input into the hybrid ML model to analyze the importance of each feature quantitatively to implement feature screening and feature interaction for prediction. This study provides new insights and has reference significance for geological feature prediction problems in the geological exploration and development process with various features, hidden feature interaction, difficult prior identification, and tough modeling. We believe that it is necessary and promising to introduce more advanced explainable feature selection and feature interactive ML algorithms into geological exploration and development.

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2. Fractal Models

Fractal theory is applicable to the study of irregular and complex natural phenomena. According to fractal geometry, the number of pores N(r) exceeding a certain pore radius r follows a power-law dependence on r (Mandelbrot, 1982):

$$N(>r) = \int_{r}^{r_{\text{max}}} P(r) dr \propto r^{-D}$$
 (1)

where r_{max} represents the maximum pore radius, and P(r) denotes the distribution density function of pore radius.

Both the volume fractal dimension (D_v) and surface fractal dimension (D_s) can quantitatively characterize the microscopic pore structure characteristics. Specifically, D_v quantitatively characterizes the space filling characteristics of the pore, while D_s quantifies the roughness of the interface between the solid and pore spaces (Dou et al., 2021; F. Wang et al., 2019; L. Wang et al., 2021). In this study, we employed six widely used fractal models, covering both classes, as shown in Table S1 in Supporting Information S1. The D of rock porous media typically ranges from 2 to 3. The D nearer to 3 corresponds to rougher pore surface, more irregular pore shape, and a higher abundance of smaller pores (Cui et al., 2022; Su et al., 2018; N. Zhou et al., 2020; H. Zhu et al., 2019).

2.1. Model I

According to Equation 1, the pore structure distribution can be expressed as:

$$\frac{dN(>r)}{dr} = ar^{-D-1} \tag{2}$$

where a is a constant.

The cumulative volume with radius below r may be calculated by:

$$V(\langle r) = \int_{r_{\min}}^{r} ar^{-D-1} Fr^{3} dr = b(r^{3-D} - r_{\min}^{3-D})$$
 (3)

where F and b are constant, the shape factor F, which equals 1 for cubic pores, quantifies pore geometry r_{\min} and r_{\max} represent the smallest and largest pore radius, respectively.

Similarly, for pores with radius above r, the total volume is given by:

$$V(>r) = \int_{r}^{r_{\text{max}}} ar^{-D-1} Fr^{3} dr = b(r_{\text{max}}^{3-D} - r^{3-D})$$
 (4)

For the wetting phase cumulative saturation $(gS_{1-S_{u}})$ with r can be calculate by:

$$S_{1-S_{Hg}} = \frac{V(\langle r)}{V} = \frac{r^{3-D} - r_{\min}^{3-D}}{r_{\max}^{3-D} - r_{\min}^{3-D}}$$
 (5)

When $r_{\min} \ll r_{\max}$, it can be simplified as

$$S_{1-S_{Hg}} = \frac{r^{3-D}}{r_{\max}^{3-D}} \tag{6}$$

According to Equation 17, it can be transformed into:

$$S_{1-S_{Hg}} = \frac{P_c^{D-3}}{P_{cd}^{D-3}} \tag{7}$$

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where P_{cd} is the displacement pressure, and P_c is the external pressure.

The *D* is determined from the slope of logarithmic plot of $S_{1-S_{Hg}}$ versus P_c (He & Hua, 1998), by D=3+ Slope *D* calculated by this model is denoted as *D1*.

2.2. Model II

The relationship between S_{Hg} and V_{Hg} for MIP can be calculated as follows:

$$S_{Hg} = \frac{V_{Hg}}{V_{\text{pore}}} \tag{8}$$

where V_{pore} represents the total pore volume of the porous medium; V_{Hg} represents the cumulative mercury volume.

Substituting Equation 8 into Equations 15 and 1, we obtain (K. Li, 2010; Shen et al., 1995):

$$S_{Hg} \propto P_c^{D-2} \tag{9}$$

Therefore, the *D* is calculated by the slope of $log(S_{Hg})$ versus $log(P_c)$, by D=2+Slop. D computed using this model is named as D2.

2.3. Model III

 $D_{\rm s}$ can be obtained from the pore volume and pore radius according to (Friesen & Mikula, 1987):

$$\operatorname{Log}\left(\frac{-dS_{Hg}}{dr}\right) \sim (2 - D)\log(r) \tag{10}$$

Combining Equation 1, the inverse relationship between r and Pc, Equation 10 can be simplified as:

$$\log\left(\frac{dS_{Hg}}{dP_c}\right) \sim (D-4)\log(P_c) \tag{11}$$

The D of the porous medium surface is computed from the slope of $\log(dS_{Hg}/dP_c)$ versus $\log(P)_c$ of which D=4+ Slop. D calculated by this model is denoted as D3.

2.4. Model IV

B. Zhang and Li (1995) developed a fractal model grounded in the energy balance between surface energy increase during nonwetting phase displacement and external work required for mercury intrusion. The surface dimension can be determined by:

$$\ln\left(\frac{W_n}{r_n^2}\right) = C + D\ln\left(\frac{V_n^{\frac{1}{3}}}{r_n}\right) \tag{12}$$

where C is a constant; Wn is the intruded work, calculated by $W_n = \sum_{i=1}^n P_i \Delta V_i$ with P_i and V_i correspond to the external pressure and intruded mercury volume at each step i, respectively; V_n represents the cumulative mercury volume intruded into pores of radius r. D derived from the slope of $\ln(W_n/r_n^2)$ versus $\ln(V_n^{\frac{1}{3}}/r_n)$, by D = Slop D calculated by this model is denoted as D4.

2.5. Model V

Neimark (1990) introduced a method to calculated D_s using the thermodynamic relationship between interfacial surface area and adsorbed mass, which does not use any adsorption model on fractal surfaces. We can calculate

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the solid/liquid interface area S as a function of Pc from the Rootare-Prenzlow equation (Rootare & Prenzlow, 1967). By combining with Equation 17, the D calculation proceeds as Equation 14:

$$S(P_c) = \left(\frac{1}{\gamma \cos \theta}\right) \int_0^V P_c dV \tag{13}$$

$$D = 2 + \frac{d[\log S(P_c)]}{d\log P_c} \tag{14}$$

where γ represents the liquid surface tension; θ denotes the wetting angle, and V corresponds to the cumulative intruded mercury volume. The D is determined from the slope of $\log S(P_c)$ versus $\log P_c$, and $D=2+\operatorname{Slop}$. D calculated by this model is denoted as D5.

2.6. Model VI

Assuming pores as a capillary bundle model (cylindrical pores with length l and radius r), N(r) is calculated using Equation 15 (Shen & Li, 1994):

$$N(r) = \frac{V_{Hg}(r)}{\pi r^2 l} \tag{15}$$

where $V_{Hg}(r)$ denotes the cumulative mercury volume with pore radius r; N(r) is the number of pores modeled as cylindrical capillaries with radius r, following the distribution described by Equation 1. The D is calculated by the incline of $\log(N(r))$ versus $\log(r)$ by D = -Slop. This model-derived dimension is designated as D6.

2.7. Multifractal

Multifractal can quantitatively describe the distribution of a physical quantity within an irregular fractal space/set and represents a heterogeneous collection of subsets with varying fractal behaviors (Paz Ferreiro & Vidal Vázquez, 2010). The strong heterogeneity of rock porous media may lead to differences of D in different pore size segments. Using the segmental regression method based on pore size distribution (PSD) to calculate D_i can better reflect the partial feature (P. Li et al., 2017; J. Wu et al., 2019). And the total D of the whole PSD can be calculated using Equation 16:

$$D = \sum_{i=1}^{n} \frac{\varphi_i}{\varphi_{\text{total}}} D_i \tag{16}$$

here, n is the number of fractal intervals divided according to the PSD; φ_i denotes the porosity (%) of various pore size ranges; D_i defines the fractal dimension of the corresponding interval, ranging from 2 to 3, as discussed above; φ_{total} is the total porosity of the rock. In this study, this method is specifically employed for samples exhibiting multiple fractals. We calculated the total fractal dimension and discussed its relation to the petrophysical properties and pore structure parameters of tight sandstone. This approach ensures completeness and accuracy in calculating D.

3. Materials and Methods

3.1. Sampling

The Kuqa Depression, a Mesozoic-Cenozoic superimposed foreland basin, is located in the northern Tarim Basin near the South Tianshan foothills (X. Guo et al., 2016). The Kelasu structural belt is the first row of thrust belts in the South Tianshan piedmont. To the west lies the Bozi-Dabei area, an imbricated structural belt formed under the strong thrust and nappe, covering an area of 2,070 km² (Tian et al., 2020), with relatively limited exploration so far. Achieving breakthroughs in natural gas exploration and development in this area is essential for ensuring long-term stable gas supply for the China West-East Gas Pipeline Project. All samples were collected from the BS

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Fm. in Bozi-Dabei area. The BS Fm. is the main producing formation, lying at a significant depth of 5,000–8,000 m (details in Figure S1 in Supporting Information S1).

3.2. Experiments

For all test samples, chloroform and distilled water were used to remove residual bitumen and salt and finally dried in a drying cabinet at 105°C for 24 hr. The irregular cutting blocks, excised from both ends of the core, were utilized for microscopic observation and X-ray diffraction analysis (XRD). The regular plunger was used for physical property analysis, and subsequently, a 1 cm section of the regular plunger excised for the HPMI experiment.

For most solid interfaces, mercury exhibits a non-wetting behavior. Thus, Pc (external pressure) is required to penetrate hydrophilic rock pores. The attraction of mercury to the surface of porous media is equal to the circumference (supposing a circular pore). The equilibrium, when external force matches resistance, can be described as Washburn (1921):

$$-\pi \mathrm{D}\gamma \cos \theta = \frac{\pi \mathrm{D}^2 P_c}{4} \tag{17}$$

Simplified as:

$$W = -\frac{4\gamma\cos\theta}{P_c} \tag{18}$$

where γ is 485 dyne cm⁻¹ for air/mercury system; θ is 140° for mercury/rock system; W is the pore diameter. The instrument operates at a maximum mercury injection pressure of 200 MPa, enabling measurement of pore radius down to 0.004 μ m. Due to the low compressible of tight sandstone, which contains more rigid components compared to shale and coal, this study omits compressibility corrections.

Pcd is defined as the mercury pressure at the intersection of the tangent to the inflection point of the mercury curve and the *y*-axis. The volume of mercury injection into pores larger than the pore size corresponding to *Pcd* is regarded as rock surface volume and can be ignored. Pittman (1992) employed the hyperbolic graphical method, utilizing the vertex as an effective indicator, while Winland (Kolodzie, 1980) defined the pore radius *R35*, at 35% mercury saturation, to evaluate reservoir physical properties. Detailed illustrations are provided in Figure S2 in Supporting Information S1.

3.3. Algorithms

3.3.1. Permutation Importance

There are many feature parameters affecting permeability, so convincing feature selection to obtain more important features is crucial to enhance the prediction model (Chandrashekar & Sahin, 2014; Guyon & Elisseeff, 2003; Kuhn & Johnson, 2018). In recent years, the PI algorithm has garnered significant attention as a model-agnostic feature selection approach, outlined in Equation 19 (Altmann et al., 2010; Fisher et al., 2018). This algorithm evaluates feature importance through randomly permuting feature values and monitoring the impact on model performance. It boasts strong interpretability and generous applicability, making it suitable for structured, small-sample data and any model. Specifically, for each feature, the algorithm generates a randomly permuted version and replaces the original feature with this permuted one, observing changes in model performance. A significant decrease in model performance indicates the importance of that feature for prediction. Compared to other commonly used feature selection algorithms, the PI algorithm solely relies on validation results and requires no model modifications, offering greater flexibility and convenience. The adoption of the PI algorithm for feature selection enables a more comprehensive and precise evaluation of feature importance, providing valuable support for subsequent model training and prediction.

$$i_j = s - 1/K \sum_{k=1}^{K} s_{k,j} \tag{19}$$

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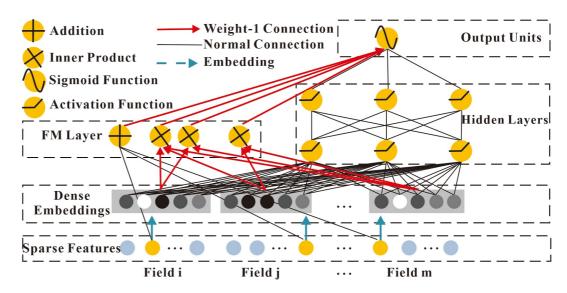


Figure 1. Wide & deep architecture of DeepFM. Both components share identical raw feature inputs, enabling simultaneous learning of low- and high-order feature interaction information directly from the original features (H. Guo et al., 2017).

where j is the feature index in the feature tabular data set A. The importance of feature j, denotes as i_j , is computed, and s represents the compute reference score of the validation model m on data A. For each repetition k in $1 \dots K$, column j of data set A is randomly shuffled to create a corrupted version, designated as $\tilde{D}_{k,j}$, and $s_{k,j}$ is the compute score of model m on corrupted data $\tilde{D}_{k,j}$.

3.3.2. DeepFM

The DeepFM model is a neural network framework that integrates deep neural networks with factorization machines (H. Guo et al., 2017). It's built upon the Wide & Deep structure, using FM to replace the LR of the Wide module, avoiding the complex feature engineering of manual construction. FM extracts features by the inner product of hidden variables of each dimension feature, having strong interpretability. The FM component in DeepFM simultaneously models linear (order-1) and pairwise (order-2) feature interactions (Rendle, 2010). The deep component, a feed-forward neural network, is designed for mining high-order feature interaction information. Notably, the FM and Deep modules share a common feature embedding part, enabling mutual optimization of low- and high-order feature representations through backpropagation (Figure 1 and Table 1). This joint training strategy enhances representation learning precision by synergizing both interaction types. And the structures of FM, Deep and embedding layer show in Figures S3, S4, and S5 in Supporting Information S1, respectively. Compared to other deep models like FNN, PNN and Wide & Deep, DeepFM can more effectively capture low- and high-order feature interactions simultaneously without the requirement for pre-training or manual feature engineering (Cheng et al., 2016; H. Guo et al., 2017; Rendle, 2010). The prediction function of DeepFM is:

$$\hat{y} = \text{sigmoid}(y_{\text{FM}} + y_{\text{DNN}}) \tag{20}$$

| Table 1 Comparison of the Models for Prediction | | | | | | | |
|---|------------------|--------------------|--------------|-------------------------------|--|--|--|
| Model | Linear (order-1) | Pairwise (order-2) | High-order | No manual feature engineering | | | |
| Classical models | $\sqrt{}$ | × | × | × | | | |
| FM | $\sqrt{}$ | \checkmark | × | × | | | |
| DNN | × | × | \checkmark | × | | | |
| DeepFM | $\sqrt{}$ | $\sqrt{}$ | $\sqrt{}$ | $\sqrt{}$ | | | |

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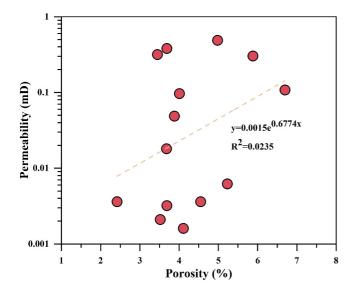


Figure 2. Porosity shows a poor correlation with permeability.

where \hat{y} is the prediction permeability, y_{FM} is the output of the FM component, and y_{DNN} is the output of the deep component, with a sigmoid function applied for prediction.

FM Model Equation: The equation for a factorization machine with a degree of d = 2 is defined as:

$$y_{\text{FM}} = w_0 + \sum_{i=1}^{n} w_i x_i + \sum_{i=1}^{n} \sum_{j=i+1}^{n} \langle v_i, v_j \rangle x_i x_j$$
 (21)

The parameters of the model that must be assessed are:

$$w_0 \in R, w \in R, V \in R^{n \times k} \tag{22}$$

And $\langle \cdot, \cdot \rangle$ denotes the dot product of two vectors of size k:

$$\langle v_i, v_j \rangle = \sum_{f=1}^k v_{i,f} \cdot v_{j,f}$$
 (23)

A row v_i within V describes the *i*th variable with k factors. $k \in \mathbb{N}_0^+$ is a hyperparameter that determines the dimensionality of the factorization. w_0

represents the global bias; w_i models the strength of the *i*th variable. $\langle v_i, v_j \rangle$ represents the inner product of eigenvector v_i and eigenvector v_j ; Vi is the hidden factor vector of x_i feature; v_j is the hidden factor vector of x_j feature. The relationship between features is represented by the dimension of the latent factor vector, which is essentially an embedding representation of the features. According to Cholesky decomposition, for any symmetric positive semi-definite matrix W, as long as k is large enough, there must be a matrix V such that $W = V \cdot V^T$ can approximate any second-order cross feature in this way. Thus, FM more effectively learns feature interactions that are rare or absent in the training data and captures order-2 feature interaction information better than former methods.

4. Results

4.1. Petrophysical Properties

The petrophysical parameters indicate that the BS Fm. is a typical tight sandstone reservoir, with porosity (φ) and permeability (K) values ranging from 2.42% to 6.70% (average of 4.27%) and 0.0016 mD to 0.486 mD (average of 0.13 mD), respectively. The porosity and permeability are weakly correlated, with the permeability varying by two orders of magnitude among core samples with similar porosities (Figure 2), indicating that the porosity alone impacts the permeability to a limited extent (Cai et al., 2019).

The reservoir quality index (RQI) and the flow zone indicator (FZI) were computed via Equations 24 and 25, respectively (Abbaszadeh et al., 1996; Amaefule et al., 1993), with detailed results in Table S2 in Supporting Information S1. The RQI and FZI are key parameters in hydraulic flow unit theory and can serve to classify highly heterogeneous reservoirs. The larger their values are, the better the reservoir quality (P. Zhao et al., 2017).

$$RQI = 0.0314 \times \sqrt{\frac{K}{\varphi}}$$
 (24)

$$FZI = RQI \times \frac{100 - \varphi}{\varphi}$$
 (25)

4.2. Pore Space Characteristics

4.2.1. Pore Types

The casting thin section and SEM results (as shown in Figure S6 in Supporting Information S1) indicate that the tight sandstone contained various types and multi-scale pores, such as residual primary intergranular pores,

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dissolution pores, intercrystalline pores and microfractures. The dominant pore types are residual intergranular and intergranular dissolution pores. The residual intergranular pores exhibited simple geometric shapes, typically triangular, polygonal or long strip, with radius generally exceeding 10 µm. They develop between skeleton grains like quartz and feldspar grains and are often filled with secondary authigenic quartz, authigenic feldspar and I/S mixed layer, resulting in relatively isolated pores with poor connectivity. Intragranular dissolution pores, formed by partial dissolution in feldspar and debris grains, are usually irregular polygons and relatively small in size, whereas the mold pores formed by complete dissolution can reach tens of microns. Intergranular dissolution pores, resulting from the dissolution of feldspar and debris, exhibit complex, irregular and varied geometries. Their size depends on the dissolution intensity, and the connectivity of the dissolution pores is relatively good. Intercrystalline pores, occurring within authigenic clay minerals such as illite and I/S mixed layers, range from several to hundreds of nanometers in size, with geometric shape determined by the mineral type, distribution and arrangement. In general, the connectivity between the pores on the plane is poor, and connected by sheet-shaped and curved-sheet-shaped throats. Various and multistage microfractures are present in some samples, including structural fractures formed by tectonic stress, grain breaking cracks and grain edge dissolution cracks. Although they contribute little to pore space, unfilled microfractures can significantly enhance permeability. Owing to the strong heterogeneity of tight sandstone, thin-section analysis alone provides limited insight into pore types and space characteristics, but cannot fully capture three-dimensional pore characteristics and real PSDs. Additionally, the pore types and distributions observed on the plane under the microscope also demonstrate strong heterogeneity. Thus, while thin-section analysis offers valuable reference, it cannot serve as the sole standard to judge the impact of pore type and structure on physical properties.

4.2.2. Pore Structure

Table S2 in Supporting Information S1 shows the HPMI experimental results. The *Pcd* shows relatively high values varying between 0.14 and 5.51 MPa (average of 1.43 MPa). The median radius (*R50*), defined as the radius at 50% mercury injection saturation, varies from 0.006 to 0.67 μm (average 0.22 μm). Sorting coefficient (*Sp*) ranges from 1.76 to 3.23 (average 2.63). In addition, the samples have relatively high maximum mercury saturation (*Smax*) (40.41%–96.96%/70.92%), but low mercury removal efficiency (*We*) (28.05%–58.54%/42.13%), and a large amount of mercury remained. These indicate poor sorting, distinct pore-throat distribution, and strongly heterogeneous and complex pore structure. The pore-throat distribution curve of tight sandstone obtained by HPMI analysis is shown in Figure S7 in Supporting Information S1. There are two distribution types, bimodal (Figure S7b in Supporting Information S1) and bias to unimodal (Figure S7c in Supporting Information S1), with maximum pore size below 2 μm. Furthermore, the macroscopic petrophysical properties like porosity, permeability, RQI and FZI show poor correlation with the parameters derived from HPMI (Figure 3). Therefore, a more effective approach for establishing tight reservoir classification and evaluation criteria involves the integrated analysis of multiple factors, including PSD, pore type and pore geometry characteristics.

Type I reservoir exhibits permeability greater than 0.01 mD, RQI greater than 0.02 (average 0.06), and FZI greater than 0.55 (average 1.43), indicating relatively good macroscopic physical properties. Most sample platform areas are narrow with relatively large slopes. It has relatively high Pcd (average 2.12 MPa), relatively low Smax (average 59.61%), and small R50 and average pore radius (Rave) (averages of 0.038 and 0.18 μ m, respectively). The PSD curve is bimodal, with peaks between 0.003 and 0.008 μ m and 0.08 and 0.7 μ m. The development of dissolution pores, including intragranular and intergranular types with complex geometric shapes, as well as grain edge dissolution fractures, observed under microscope. Some samples develop fractures running through thin section, including open, half-open or closed fractures.

Type II reservoir has permeability less than 0.01 mD, RQI below 0.012 (average 0.009), and FZI under 0.49 (average 0.24), suggesting poor macroscopic physical properties. Most sample platform areas are relatively wider with gentler slopes. It features relatively lower Pcd (average 0.52 MPa), higher Smax (average 86.01%), and larger R50 and Rave (averages of 0.34 and 0.47 μ m, respectively). The PSD curve bias to unimodal with peak distribution ranging from 0.2 to 1 μ m. Under the microscope, intergranular pores with regular geometric shapes are relatively developed, and most of them exist relatively isolated on the plane or are connected by sheet-like throats.

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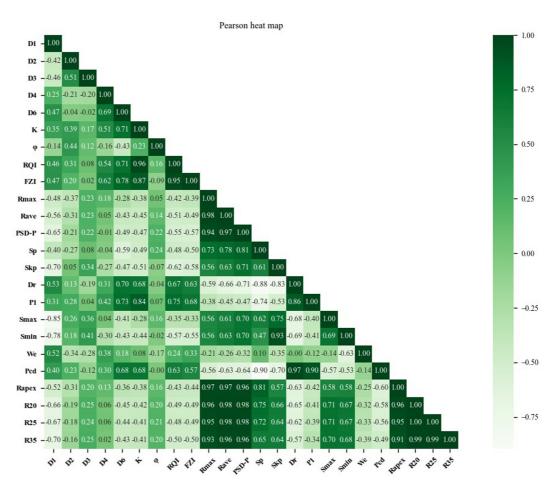


Figure 3. Heatmap reflecting the relationship among fractal dimension, HPMI parameters and macro petrophysical parameters. Note that PSD-P is the radius of pore throat distribution peak; Dr represents the relative sorting coefficient; PI is the characteristic structure coefficient; Smin represents the final remaining mercury saturation. Notes. The Pearson correlation coefficient (|R|) are categorized as: no (|R| < 0.01), negligible ($0.01 \le |R| < 0.2$), weak ($0.2 \le |R| < 0.3$), moderate ($0.3 \le |R| < 0.4$), strong ($0.4 \le |R| < 0.7$), and very strong ($0.7 \le |R| \le 1$) association (L. Wang & Fu, 2018).

4.3. Fractal Dimension

The D calculations utilize data points above P_{cd} . Table S3 in Supporting Information S1 and Figures S8, S9, and S10 in Supporting Information S1 detail the results from the above six fractal models.

4.3.1. Two-Regions Fractal Models

For most samples, two fractal regions can be identified via model I. Region I, corresponding to large pores, exhibits *D* ranging from 2.59 to 2.93. Region II, corresponding to small pores, shows *D* between 1.15 and 2.95, with some not having physical meaning. The total fractal dimension *D1* distribution spans 2.58 to 2.89.

Model II also reveals two distinct regions. For region I, the calculated D ranges from 2.39 to 3.92, with some not having physical meaning. For region II, the calculated D ranges from 2.09 to 2.29. The total fractal dimension D2 is varied between 2.12 and 2.38.

Model III also yields two fractal regions. For region I, the calculated D is between 1.70 and 2.94, whereas for region II, D ranges from 2.64 to 4.13. Some samples, such as S6 and S2, have no fractal geometric meaning for the large or small pores. The total fractal dimension D3 ranges between 2.01 and 2.91.

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Fractal models IV, V and VI are all satisfied in the whole range of mercury injection characterization apertures, and all samples get only one fractal dimension, *D4* and *D6* ranges from 2.55 to 2.84, and 2.13 to 2.38, respectively, and have real physical meaning. However, the *D5* of all samples are greater than 3 (ranging from 3.06 to 3.20), thus, the fractal results have no physical meaning.

5. Discussions

5.1. Differences Among Fractal Models

All D5 values are greater than 3, indicating that the fractal results have no real physical meaning, so it is deemed inapplicable and excluded from further discussion. Both models IV and VI recognize only one fractal region, D4 and D6 can represent the entire pore size range measured by HPMI. And they are significantly positively correlated (Figure 3), indicating their consistency in characterizing reservoir pore structure heterogeneity. Furthermore, basically all samples exhibit obvious inflection points, making fractal models I, II, and III effective for distinguishing different fractal regions, which is consistent with former research (Cui et al., 2022; Dou et al., 2021). We calculate D of Regions I and II, distinguished through segmental regression applied to PSD and obtain the total value using Equation 16. As shown in Figure 3, correlation analysis reveals that D3 is positively correlated with D2, whereas D1 is negatively correlated with D2. This indicates that using D derived from model I, model II or III to analyze rock pore structure and physical properties may lead to opposite conclusions. A higher D6 value signifies a complex pore structure, lower porosity, higher permeability, and elevated FZI and RQI values; in addition, the correlation between D6 and physical properties is the best. In contrast, D1, D2, and D3 show no significant or only weak correlations with the macroscopic physical parameters. In summary, D6 can effectively describe the fractal characteristics across the entire PSD and shows the best correlation with reservoir physical properties. Therefore, it is the optimal choice for evaluating tight sandstone reservoirs physical properties.

5.2. Factors Influencing Permeability

Permeability is the comprehension reflect of multiple geological factors. It is necessary to deeply analyze the factors influencing permeability before modeling (R. Guo et al., 2020; Oluwadebi et al., 2019). The results of correlation analysis of the HPMI parameters, fractal dimension, main mineral types, and permeability and porosity, are shown in Figures 3 and 4: (a) Permeability is correlated with certain HPMI experimental parameters. For example, Rave and Skp are negatively correlated with permeability, indicating that the more the PSD is biased toward smaller pore sizes, the higher the permeability; while P1 is positively correlated with permeability, reflecting the better the degree of pore-throat sorting and pore-throat connectivity, the better the pore structure and the higher the permeability, that is, permeability is affected by pore size and distribution (Cai et al., 2019; F. Wang et al., 2019). (b) D6 is positively correlated with permeability for type I reservoir and negatively correlated for type II reservoir, with a similar pattern observed for porosity. For type I reservoir, the presence of dissolution pores with complex geometric but relatively better connectivity and fractures leads to a higher D and consequently higher permeability, although the porosity inclines to decrease. Conversely, for type II reservoirs, which are dominated by intergranular pores with relatively poor connectivity, a higher D indicates a more complex pore structure, resulting in lower permeability and porosity (Lai et al., 2016). (c) The type and genetic mechanism of minerals affect pore characteristics (e.g., type, geometry and size distribution), thereby affecting reservoir physical properties. Researchers have investigated the relationships among mineral composition, microscopic pore structure and fractal dimension (Al-Kharra'a et al., 2023; Qiao et al., 2022; S. Zhao et al., 2022), demonstrating that distinct mineral compositions exert varying controls on these parameters, so we will not elaborate here. Minerals and physical properties exhibit complex and diverse relationships, even with opposing trends observed in different reservoirs types (Figure 4). In Type I reservoirs, an increase in carbonate mineral indicates enhanced dissolution and the potential presence of microfractures filled with carbonate mineral, leading to increased permeability and porosity. In contrast, higher quartz mineral and clay mineral contents suggest relatively weaker dissolution and poorer pore connectivity, resulting in lower permeability and a tendency for the porosity to decrease. For Type II reservoirs dominated by intergranular pores, carbonate mineral primarily blocks pores and throats, belong to destructive diagenesis. Feldspar and quartz, which support intergranular pores

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1944/973, 2025, 7, Downloaded from https://agupubs.onlinelibrary.wiley.com/doi/10.1029/2024WR03379 by China University Of Petroleum, Wiley Online Library on [21.07/2025]. See the Terms and Conditions (https://onlinelibrary.wiley.com/doi/10.1029/2024WR03379 by China University Of Petroleum, Wiley Online Library on [21.07/2025]. See the Terms and Conditions (https://onlinelibrary.wiley.com/doi/10.1029/2024WR03379 by China University Of Petroleum, Wiley Online Library on [21.07/2025]. See the Terms and Conditions (https://onlinelibrary.wiley.com/doi/10.1029/2024WR03379 by China University Of Petroleum, Wiley Online Library on [21.07/2025]. See the Terms and Conditions (https://onlinelibrary.wiley.com/doi/10.1029/2024WR03379 by China University Of Petroleum, Wiley Online Library on [21.07/2025].

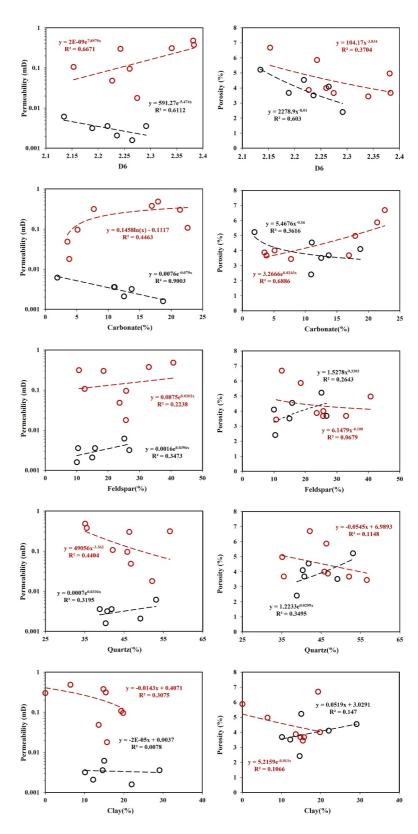


Figure 4. Correlation analysis of permeability, porosity with D6 and different minerals contents.

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Table 2
The Results and Errors of Classic Permeability Prediction Models

| Model | Equation | MSE (mD ²) | RMSE (mD) |
|---------------|--|------------------------|-----------|
| Pittman | $logK = 3.4699log\phi - 1.3524logRapex - 4.3405$ | 0.425 | 0.652 |
| Winland(R10) | $logK = 3.7789log\phi - 1.3399logR10 - 4.1792$ | 0.375 | 0.612 |
| Winland(R15) | $logK = 3.9089log\phi - 1.3444logR15 - 4.3974$ | 0.347 | 0.589 |
| Winland(R20) | $logK = 4.0795log\phi - 1.3208logR20 - 4.6299$ | 0.326 | 0.571 |
| Winland(R25) | $logK = 4.2632log\phi - 1.2164logR25 - 4.8432$ | 0.318 | 0.564 |
| Winland(R30) | $logK = 4.5027log\phi - 1.0021logR30 - 5.0482$ | 0.320 | 0.566 |
| Winland (R35) | $logK = 5.0402log\phi$ 0.8584logR355.4809 | 0.312 | 0.559 |
| Winland(R40) | $logK = 5.1349log\phi - 0.6969logR40 - 5.5627$ | 0.338 | 0.581 |

Note. The bold values represent the results of the optimal model.

development, positively impact permeability. Consequently, as the carbonate mineral content decreases and the quartz and feldspar contents increase, permeability becomes higher and porosity tends to increase.

In summary, the factors affecting permeability are highly diverse and complexly interactive, making it challenging to characterize permeability fully and well through one or a few fixed parameters. Even though the interplay between factors influencing permeability has been meticulously analyzed and discussed, manual modeling remains challenging. Therefore, it is essential to employ feature interaction algorithms based on ML to automatically mine hidden geological information for more precise and intelligent permeability prediction modeling.

5.3. Winland and Pittman Permeability Prediction Models

Percolation theory establishes characteristic length as the dominant control on seepage properties (Katz & Thompson, 1986, 1987). It can be calculated through HPMI experiment (Kolodzie, 1980; Pittman, 1992). Pittman (1992) employed mathematical methods to confirm *Rapex* as the characteristic length controlling permeability, while Kolodzie (1980) proposed that the characteristic length matches the pore radius (*R35*) at 35% mercury saturation and provided empirical formulas. We obtained the pore radius at various mercury saturation levels of 10%, 15%, 20%, 25%, 30%, 35%, and 40% and denoted these values as *R10*, *R15*, *R20*, *R25*, *R30*, *R35*, and *R40*, respectively. On the basis of classic empirical formulas, the permeability is predicted through the multiple regression method based on the least square algorithm. The accuracy of each model is evaluated using the mean squared error (MSE) and the root mean squared error (RMSE) (Table 2). However, since the dimension of the RMSE is consistent with that of the predicted value, making it more intuitively interpretable and widely applicable in practical scenarios.

MSE =
$$\frac{\sum_{i=1}^{n} (K_{\text{exp}} - K_{\text{pred}})^2}{n}$$
 (26)

$$RMSE = \sqrt{\frac{\sum_{i=1}^{n} (K_{exp} - K_{pred})^2}{n}}$$
 (27)

where K_{exp} is the experimentally measured permeability, and K_{pred} is the permeability predicted by the model.

The MSE is in a range from 0.312 to 0.425, the RMSE ranges from 0.559 to 0.652, and the Winland (*R35*) model has the smallest error among them. Possible explanations for the unsatisfactory prediction performance of the Pittman and Winland models are as follows: (a) As an important parameter of these models, the relationship between porosity and permeability is not clear in the study area. (b) A single characteristic length parameter cannot fully characterize the pore structure characteristics, and more parameters need to be considered for permeability prediction.

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| (a) | Weight | Feature | (b) | Weight | Feature |
|---|--|---|--------------------------------------|--|---|
| 0.4005 | ± 0.0986 | D6 | 2.169 | 98 ± 1.2837 | D6 |
| | ± 0.2906 | P1 | 0.31 | 11 ± 0.3549 | Carbonate |
| 0.0524 | ± 0.0076 | φ | 0.044 | 19 ± 0.0449 | Dr |
| 0.0355 | ± 0.0797 | Dr | | 0 ± 0.0000 | R20 |
| 0.0175 | ± 0.0146 | Carbonate | | 0 ± 0.0000 | φ |
| 0.0156 | ± 0.0178 | Clay | | 0 ± 0.0000 | Rmax |
| 0.0074 | ± 0.0139 | Feldspar | | 0 ± 0.0000 | Rave |
| 0.0071 | ± 0.0138 | Pcd | | 0 ± 0.0000 | PSD-P |
| 0.0031 | ± 0.0045 | Sp | | 0 ± 0.0000 | Skp |
| 0.0028 | ± 0.0343 | R25 | | 0 ± 0.0000 | R25 |
| 0.0019 | ± 0.0032 | Rapex | | 0 ± 0.0000 | Smin |
| 0.0014 | ± 0.0085 | R20 | | 0 ± 0.0000 | Pcd |
| 0.0002 | ± 0.0039 | Rave | | 0 ± 0.0000 | Quartz |
| 0.0001 | ± 0.0061 | Quartz | | 0 ± 0.0000 | Clay |
| -0.0127 | ± 0.0767 | Skp | | 0 ± 0.0000 | Feldspar |
| -0.0209 | ± 0.0318 | PSD-P | | 0 ± 0.0000 | Rapex |
| -0.0217 | ± 0.0726 | Smin | | 0 ± 0.0000 | P1 |
| -0.0228 | ± 0.0255 | Rmax | -0.025 | 51 ± 0.0615 | Sp |
| | | | | | |
| (c) | Weight | Feature | (d) | Weight | Feature |
| 0.6672 | 2 ± 1.6621 | D6 | 2.039 | 96 ± 0.7039 | Feature D6 |
| 0.6672 0.4944 | 2 ± 1.6621 4 ± 1.9970 | D6 Skp | 2.039 | 96 ± 0.7039 38 ± 0.0006 | D6 φ |
| 0.6672 0.4944 0.3455 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 | D6 | 2.039 | 96 ± 0.7039 88 ± 0.0006 0 ± 0.0000 | D6 φ Smin |
| 0.6672 0.4944 0.3455 0.1559 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 | D6 Skp Clay φ | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave |
| 0.6672 0.4944 0.3455 0.1559 0.1266 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 | D6 Skp Clay φ Carbonate | 2.039 | 96 ± 0.7039 98 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 | D6 Skp Clay φ Carbonate Sp | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 3 ± 0.0960 2 ± 0.0653 0 ± 0.0000 | D6 Skp Clay φ Carbonate Sp Smin | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 | D6 Skp Clay Φ Carbonate Sp Smin R20 | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd |
| 0.6672 0.4944 0.3455 0.1558 0.1266 0.0532 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz |
| 0.6672 0.4944 0.3455 0.1559 0.0532 (() | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 1 ± 0.0000 1 ± 0.0002 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 PSD-P | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz Clay |
| 0.6672 0.4944 0.3455 0.1559 0.0532 (() () () () () () () () () () () () () | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 1 ± 0.0000 1 ± 0.0002 0 ± 0.0068 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 PSD-P Dr | 2.039 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz Clay Carbonate |
| 0.6672 0.4944 0.3455 0.1559 0.0532 (() () () () () () () () () () () () () | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 1 ± 0.0002 0 ± 0.0068 5 ± 0.0131 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 PSD-P Dr Rave | 2.038 0.003 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz Clay Carbonate Feldspar |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 (() () () () () () () () () () () () () | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 1 ± 0.0002 0 ± 0.0068 5 ± 0.0131 5 ± 0.0231 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 PSD-P Dr Rave Quartz | -0.00° | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 15 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz Clay Carbonate Feldspar Rapex |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 (() () () () () () () () () () () () () | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 8 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 1 ± 0.0002 0 ± 0.0068 5 ± 0.0131 6 ± 0.0454 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 PSD-P Dr Rave Quartz Rmax | -0.003 -0.003 | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 15 ± 0.0000 15 ± 0.0021 30 ± 0.0047 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz Clay Carbonate Feldspar Rapex Rmax |
| 0.6672 0.4944 0.3455 0.1559 0.1266 0.0532 (0 0.0001 -0.0004 -0.0065 -0.0165 -0.0174 | 2 ± 1.6621 4 ± 1.9970 5 ± 0.4146 9 ± 0.0416 6 ± 0.0960 2 ± 0.0653 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 1 ± 0.0002 0 ± 0.0068 5 ± 0.0131 5 ± 0.0231 | D6 Skp Clay φ Carbonate Sp Smin R20 R25 Pcd P1 PSD-P Dr Rave Quartz | -0.00° -0.00° -0.00° -0.00° | 96 ± 0.7039 38 ± 0.0006 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 0 ± 0.0000 15 ± 0.0000 | D6 φ Smin Rave PSD-P Sp Dr R20 R25 Pcd Quartz Clay Carbonate Feldspar Rapex |

Figure 5. The results of PI using different validation models (a) RF; (b) DT; (c) LightGBM; (d) XGBoost. The initial value in each row quantifies model performance attenuation. The values behind \pm denotes the standard deviation across multiple tests. Negative values signify improved prediction accuracy when features are perturbed than the real data.

5.4. Explainable Hybrid ML Model

5.4.1. Feature Selection

Generally, features that are poorly correlated with the target variable will increase the prediction model error. Therefore, we first use the Pearson algorithm to conduct correlation analysis on numerous HPMI parameters, mineral contents, and fractal dimensions, combined with expert knowledge, and 18 features are screened out. Then we utilize the PI algorithm with four different validation models, RF, DT, LightGBM and XGBoost algorithms to evaluate the permutation feature importance, thereby achieving feature selection and dimensionality reduction (Figure 5). The RF and LightGBM validation models are the optimal validation algorithms for feature selection due to the consistency of the results and consistent with the geological cognition. Thus, we yield two optimal feature sets for prediction: D6, P1, φ , Dr, Carbonate, Clay, Feldspar, Pcd, Sp, R25 (10 features) and D6,

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Table 3 *Errors of the Permeability Prediction Model*

| Model | MSE (mD ²) | RMSE (mD) | Error reduction percent |
|---------------------------|------------------------|-----------|-------------------------|
| Winland (R35) | 0.312 | 0.559 | - |
| Linear (6 Features) | 0.262 | 0.451 | ↓19.32% |
| SVR (6 Features) | 0.256 | 0.500 | ↓10.55% |
| XGBoost (6 Features) | 0.294 | 0.474 | ↓15.21% |
| LightGBM (6 Features) | 0.319 | 0.488 | ↓12.70% |
| Radom Forest (6 Features) | 0.245 | 0.448 | ↓19.86% |
| FM (6 Features) | 0.288 | 0.536 | ↓4.11% |
| DNN (6 Features) | About 100 | About 10 | - |
| DeepFM (18 Features) | 0.397 | 0.542 | ↓3.04% |
| PI-DeepFM (10 Features) | 0.334 | 0.475 | ↓15.03% |
| PI-DeepFM (6 Features) | 0.240 | 0.415 | ↓25.76% |

Note. Error reduction percent: The percentage reduction of error (RMSE) between the three DeepFM models and the best classical Winland (*R*35) model. Unsatisfactory DNN. The bold values represent the results of the optimal model.

Skp, Clay, φ , Carbonate Sp (6 features). Notably, consistently ranked as the most influential parameter across the four validation models, with DT and XGBoost attributing importance results even greater than 1-a statistical anomaly underscoring its dominance to a certain extent. This aligns with the above analysis.

5.4.2. Prediction Model

Considering that there may be low- and high-order feature interaction and to assess the rationality and reliability of feature selection again, we apply the DeepFM algorithm using features after and before PI feature selection. A total of 50 rounds of training were conducted, with 100 iterations per round, and strictly apply the following criteria to select the best model: the losses of both the training and test sets must decrease simultaneously, and the difference between the two should be the smallest. The error value of the best model for 50 rounds does not fluctuate notably. The arithmetic mean of error values are shown in Table 3. The DeepFM model with 18 features achieves an RMSE of 0.542, representing a 3.04% improvement over the Winland (*R35*) model. After PI feature selection, the PI-DeepFM models demonstrate a significantly enhanced performance: the RF validation version (10 Features) reduces the RMSE to 0.475 (an improvement of 15.03%), whereas the LightGBM validation version (6 Features) achieves RMSE of 0.415 (an improvement of 25.76%). This progressive enhancement with feature reduction suggests improved model stability and convergence. In addition, the PI-DeepFM model outperforms classical empirical model, Liner model, SVR model, FM model, Tree-based models, and DNN model by effectively automated mining of both low- and high- order feature interactions without manual engineering. In conclusion, these results validate the distinctive advantages of DeepFM in modeling permeability under geologically complex and interaction-intensive conditions.

This approach offers a novel ML method and idea for permeability prediction of rock porous media with complex pore structure. Moreover, most geological features often interact and influence each other; thus, it is difficult to discover potential feature interaction information and model them in detail manually. Therefore, the introduction of advanced explainable feature selection and feature interactive information mining ML algorithms is necessary for geological exploration and development, and we believe that the PI-DeepFM algorithm has significant potential in the geological field.

6. Conclusions

1. *D6* has the best correlation with petrophysical properties and can be directly calculated to obtain a single fractal dimension representing the entire pore size range. Therefore, the use Model VI to characterize the pore structure of tight sandstone is recommended. A higher *D6* indicates a more complicated pore structure and a reduction in porosity, whereas permeability increases due to the development of microfractures and dissolution pores with complex geometric shapes.

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- 2. *D6* can characterize the pore structure well. In addition, the quantitative evaluation results of the PI model indicate that the most important factor affecting permeability prediction is *D6*. Overall, the PI model results verify the geological understanding well and the model is explainable and quantitative.
- 3. We evaluate classical permeability prediction models, and the best model is the Winland (*R35*) model with an RMSE of 0.559. The RMSE of the PI-DeepFM (6 Features) model based on the PI-DeepFM algorithm is 25.76% lower than that of the Winland (*R35*) model, and obviously better than the FM and DNN models.
- 4. The novel hybrid explainable ML permeability prediction model we proposed provides a new way of thinking for geological feature prediction problems with feature interaction concealment, difficult prior identification, and challenging manual modeling. We believe that it is both necessary and promising to introduce more advanced explainable feature selection and feature interactive information mining AI algorithms into geological exploration and development.

Data Availability Statement

The data and code associated with this work are available at Huo (2024a, 2024b).

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